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EVOLUTION OF NONLINEAR WAVE GROUPS ON WATER

OF SLOWLY-VARYING DEPTH

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Contract No. DAJA37-82-C-0300

Fifth Periodic Report, September 1983 -

February   1984

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REPORT DOCUMENTATION PAGE		READ INSTRUCTIONS BEFORE COMPLETING FORM
1. REPORT NUMBER	2. GOVT ACCESSION NO. AD-A141260	3. RECIPIENT'S CATALOG NUMBER
4. TITLE *(and Subtitle)* Evolution of nonlinear wave groups on water of slowly-varying depth		5. TYPE OF REPORT & PERIOD COVERED Periodic report September 83 - February 84
		6. PERFORMING ORG. REPORT NUMBER
7. AUTHOR(s) Ruth Iusim and Michael Stiassnie		8. CONTRACT OR GRANT NUMBER(s) DAJA37-82-C-0300
9. PERFORMING ORGANIZATION NAME AND ADDRESS Department of Civil Engineering Technion - Israel Inst. of Sci & Tech. Haifa, Israel.		10. PROGRAM ELEMENT, PROJECT, TASK AREA & WORK UNIT NUMBERS 61102A-IT161102-BH57-01
11. CONTROLLING OFFICE NAME AND ADDRESS USARDSG-UK PO Box 65 FPO NY 09510		12. REPORT DATE Feb 84
		13. NUMBER OF PAGES 32
14. MONITORING AGENCY NAME & ADDRESS*(if different from Controlling Office)*		15. SECURITY CLASS. *(of this report)* Unclassified
		15a. DECLASSIFICATION/DOWNGRADING SCHEDULE

16. DISTRIBUTION STATEMENT *(of this Report)*

Approved for public release; distribution unlimited

17. DISTRIBUTION STATEMENT *(of the abstract entered in Block 20, if different from Report)*

18. SUPPLEMENTARY NOTES

19. KEY WORDS *(Continue on reverse side if necessary and identify by block number)*

Water waves; gravity waves; non-linear wave groups; Zakharov equation; shoaling waves

20. ABSTRACT *(Continue on reverse side if necessary and identify by block number)*

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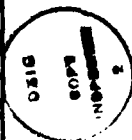
OF SLOWLY VARYING DEPTH

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Abstract

An approximate analytical solution describing the shoaling of
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about the wave-induced mean current and set down.

1. INTRODUCTION

The shoaling of weakly nonlinear surface wave groups is important
to the understanding of coastal wave climate and coastal flow regime.

In the past, most efforts concentrated on the equally important
though simpler problem of shoaling of wave-trains (i.e. monochromatic
wave groups), for details see Stiassnie & Peregrine (1980).

The first mathematical formulation for shoaling of wave-groups
was given by Djordjevic' and Redekopp (1978), and in a somewhat improved
version by Stiassnie (1983). This formulation is limited to cases
where the water depth is small compared to the group-length. Equations
suitable for water depths of the order of the group-length are
deduced in Peregrine (1983); combining the constant depth model by
Davey and Stewartson (1974) and the higher-order model for infinitely
deep water by Dysthe (1979).

The only available solutions are those for the shoaling of isolated
wave-packets (solitons), which were originally given by Djordjevic'
and Redekopp in their 1978 paper. They predict that a soliton
envelope can undergo fission only if it propagates into deeper water.
By heuristic assumptions for the evolution along the slope, they also
estimate the number of solitons emitted after a single soliton descends
from a shallower shelf. A more recent study, Turpin, Benmoussa and Mei
(1983) confirms these results qualitatively, but not quantitatively.

To the best of our knowledge, no results for shoaling of wave-groups (i.e. _modulated_ wave-trains) have been presented so far. These modulated wave-trains are of particular importance since almost every wave-train will eventually become modulated due to its intrinsic Benjamin-Feir instability. The aim of the present paper is to throw some light on the evolution during the shoaling of a modulated wave-train and its influences on the mean free surface and the wave-induced mean flow.

Sections 2, 3 and 4 outline the derivation of the mathematical model and its simplifications to a level which enables an analytical solution. The model presented at the end of section 2 is rather general; it consists of a coupled system of equations for the complex wave envelope and the induced mean flow potential. In section 3 we add the assumption of periodicity of the modulation, which leads to decoupling of the system of equations. The result is a nonlinear Schrödinger equation with a variable (depth dependent) coefficient, which can be solved by reasonable numerical efforts. In section 4 we adopt the assumption of _three-waves systems_, which has been used for constant depth in Stiassnie and Kroszynski (1982), and which together with a W.K.B. type approach enable an analytic, though asymptotic solution. This asymptotic solution is compared with numerical results in section 5. Sections 6, 7 and 8 include a detailed presentation and discussion of the physical results obtained from out calculations.

2. EVOLUTION EQUATIONS

Assuming irrotational motion, there exists a velocity potential $\phi(x,z,t)$ which satisfies Laplace's equation:

$$\phi_{xx} + \phi_{zz} = 0, \tag{2,1}$$

t is time, x is the horizontal coordinate in the direction of wave propagation, and z is the vertical coordinate pointing upward from the undisturbed free surface.

The boundary condition on the bottom, $z = -h(x)$, is

$$\phi_z = -h'(x)\phi_x. \tag{2.2}$$

The boundary conditions on the free surface, $z = \zeta(x,t)$, are the kinematic condition:

$$\phi_z = \zeta_t + \phi_x\zeta_x, \tag{2.3}$$

and the dynamic condition:

$$2g\zeta + 2\phi_t + \phi_x^2 + \phi_z^2 = 0. \tag{2.4}$$

Considering situations for which the depth h(x) as well as the wave input vary slowly, we assume all wave-field properties to be slowly changing. To render the term 'slowly' explicit we introduce a small nondimensional parameter ϵ which is a measure of the wavy surface slope, and define the following new variables:

$$\tau = \epsilon(\int^x \frac{dx}{\Omega'} - t), \tag{2.5a}$$

$$\xi = \epsilon^2 x; \tag{2.5b}$$

where $\Omega' = \partial\Omega/\partial K$ is the group velocity. $\Omega^2 = gKth(Kh)$ is the linear dispersion relation, relating the leading order constant frequency Ω to the leading order wave-number $K(\xi)$ and the water depth $h(\xi)$.

Having the wave-groups (i.e. wave fields with narrow spectra) in mind the velocity potential ϕ and the free surface displacement ζ are expanded in Fourier series:

$$\phi = \phi_o(x,z,t) + \{\phi_1(\tau,\xi,z)e^{i\theta} + \phi_2(\tau,\xi,z)e^{2i\theta} + \ldots c.c.\} \quad (2.6a)$$

$$\zeta = \zeta_o(\tau,\zeta) + \{\zeta_1(\tau,\xi)e^{i\theta} + \zeta_2(\tau,\xi)e^{2i\theta} + \ldots + c.c \} \quad (2.6b)$$

where $\theta = (\int^x K(\xi)dx - \Omega t)$, and c.c. stands for the complex conjugate. With ε chosen to be small, the functions $\phi_j(\tau,\xi,z)$ and $\zeta_j(\tau,\xi)$ for $j \geqslant 1$ are expanded formally in power series of ε :

$$\phi_j(\tau,\xi,z) = \sum_{m=j}^{\infty} \varepsilon^m \phi_{mj}(\tau,\xi, z) \quad (2.7a)$$

$$\zeta_j(\tau,\xi) = \sum_{m=j}^{\infty} \varepsilon^m \zeta_{mj}(\tau,\xi) \quad (2.7b)$$

The mean water level ζ_o, and the induced mean flow potential ϕ_o, require a special treatment and are best written as:

$$\zeta_o(\tau,\xi) = \varepsilon^2 \zeta_{20}. \quad (2.7c)$$

$$\phi_o(x,z,t) = \varepsilon\phi_{10}(\tau,\xi,z) + \varepsilon^2\phi_{20}(\xi) \cdot t. \quad (2.7d)$$

Note that it turns out that while ϕ_{mj} and ζ_{mj} for $j \geqslant 1$ are all of

order one, the order of ϕ_{10}, ζ_{20} and ϕ_{20} is between one for Kh = 0(1) and ε for Kh $\rightarrow \infty$. The term $\varepsilon^2 \phi_{20}(\xi) \cdot t$, in Eq. (2.7c), is needed to suppress terms that grow boundlessly with time at higher order, as was shown in Stiassnie (1983). Following the method of derivation used by Djordjevic' and Redekopp (1978) but using Eqs. (2.7c,d), instead of expanding ϕ_o and ζ_o in power series of ε (which is justified only for Kh = 0(1)), we obtain a system of evolution equation (2.9a,b,c,d) for ϕ_{10} and the complex wave envelope,

$$A(\tau, \xi) = -2i\varepsilon \zeta_{11} \tag{2.8}$$

as follows:

$$\frac{i}{2\Omega'} \frac{\partial \Omega'}{\partial \xi} A + iA_\xi + \frac{\Omega''}{2(\Omega')^3} A_{\tau\tau} - \frac{\varepsilon^{-2}\beta_1}{\Omega'} |A|^2 A = (\beta_2 \phi_{10_\tau} + \beta_3 \phi_{20}) \frac{A}{\Omega'};$$

$$z=0, \tag{2.9a}$$

$$\phi_{10_{zz}} + \frac{\varepsilon^2}{(\Omega')^2} \phi_{10_{\tau\tau}} = 0 ; \quad -h \leqslant z \leqslant 0, \tag{2.9b}$$

$$\phi_{10_z} + \frac{\varepsilon^2}{g} \phi_{10_{\tau\tau}} = \frac{g\beta_2}{2\Omega} (|A|^2)_\tau \quad ; \quad z = 0 \tag{2.9c}$$

$$\phi_{10_z} = 0; \qquad\qquad z=-h. \tag{2.9d}$$

The depth dependent coefficients are:

$$\beta_1 = \frac{gK^3}{2\Omega} \cdot \frac{9-12th^2(Kh)+13th^4(Kh)-2th^6(Kh)}{8th^3(Kh)} \tag{2.10a}$$

$$\beta_2 = \frac{K^2}{2\Omega} \cdot (\frac{2\Omega}{K\Omega'} + sech^2(Kh)) \tag{2.10b}$$

$$\beta_3 = \frac{K^2}{2\Omega} sech^2(Kh) \tag{2.10c}$$

A set of modulation equations equivalent to (2.9), but for constant depth, has been recently derived from the finite depth Zakharov equation, see Stiassnie and Shemer (1984). The mean water level ζ_o is given by:

$$\zeta_o = \frac{\varepsilon^2}{g} (\phi_{10_\tau} - \phi_{20}) - \frac{gK^2}{4\Omega^2} \operatorname{sech}^2(Kh) \cdot |A|^2 \qquad (2.11)$$

and is of order between ε^2 to ε^3, depending on the water depth.

3. PERIODIC CASES

Restricting the discussion to cases for which the complex wave envelope $A(\tau,\xi)$ is periodic in τ, and assuming zero averaged (over τ) mass flow in the x-direction (as in the case of an impermeable beach) enables the decoupling of the system (2.9). For these cases A is governed by the nonlinear Schrödinger equation:

$$\frac{1}{2\Omega'} \frac{\partial \Omega'}{\partial \xi} A + iA_\xi + \frac{\Omega''}{2(\Omega')^3} A_{\tau\tau} - \frac{\varepsilon^{-2}\alpha_1}{\Omega'} |A|^2 A = \frac{\varepsilon^{-2}\alpha_2}{\Omega'} \overline{|A|^2} A, \qquad (3.1)$$

where the bar denoted averaging over τ, and

$$\alpha_1 = \beta_1 + \frac{g^2\beta_2^2}{2\Omega(1-gh/(\Omega')^2)} \qquad (3.2a)$$

$$\alpha_2 = -\frac{gK\Omega'}{2h\Omega} (\beta_2+\beta_3) - \frac{g^2\beta_2^2}{2\Omega(1-gh/(\Omega')^2)} \qquad (3.2b)$$

The induced mean flow potential is given by:

$$\phi_{10} = \frac{-\varepsilon^{-2}\overline{|A|^2}gK\Omega'}{2\Omega h} \cdot \tau + \sum_{n=1}^{\infty} a_n(z)\{b_n(\xi)e^{2\pi in\tau/\gamma} + b_{-n}(\xi)e^{-2\pi in\tau/\gamma}\}$$

$$(3.3)$$

where γ is the period of $A(\tau_{\xi})$;

$$a_n = \cosh(2\pi\varepsilon n(z+h)/\Omega'\gamma), \qquad n = 1,2,\ldots; \tag{3.3a}$$

$$b_n = \frac{\varepsilon^{-2}g^2\beta_2}{2\Omega\{\frac{2\pi n g}{\varepsilon\Omega'\gamma}\sinh(\frac{2\pi n\varepsilon h}{\Omega'\gamma})-(\frac{2\pi n}{\gamma})^2\cosh(\frac{2\pi n\varepsilon h}{\Omega'\gamma})\}} \cdot C_n; \tag{3.3b}$$

and C_n are the Fourier coefficients of $(|A|^2)_{\tau}$:

$$(|A|^2)_{\tau} = \sum_{n=1}^{\infty}\{C_n(\xi)e^{2\pi i n\tau/\gamma}+C_{-n}(\xi)e^{-2\pi i n\tau/\gamma}\} \tag{3.4}$$

The convergence of the Fourier series (3.4) was assumed to be independent of ε.

The potential ϕ_{20} which is needed to calculate the mean water level ζ_o, see Eq. (2.11), is given by:

$$\phi_{20} = -\varepsilon^{-2}\overline{|A|^2}\ gK\Omega'/2\Omega h. \tag{3.5}$$

Eq. (3.1) is identical to Eq. (4.9) in Stiassnie (1983), which has been derived assuming no vertical dependence of the induced mean flow. The latter assumption is strictly valid only when $Kh = 0(1)$, namely, for cases where the water depth h is small compared to the group length $(\varepsilon K)^{-1}$. For these cases the present $\phi_{o,x}$ and ζ_o yield Eqs. (3.7) and Eq. (3.8) of Stiassnie (1983).

A simpler dimensionless form of Eq. (3.1) is obtained by means of the transformations

$$\psi = \varepsilon^{-1}\left(\frac{2\Omega^5\Omega'}{g^3}\right)^{\frac{1}{2}} A \exp i \left(\overline{|A|^2} \int_{x_\infty}^{x} \frac{\alpha_2 dx}{\Omega'}\right); \tag{3.6a}$$

$$T = \tau/\gamma; \qquad X = \frac{1}{\gamma^2} \int_{\varepsilon_\infty}^{\xi} \frac{\Omega''}{2(\Omega')^3} d\xi. \tag{3.6b}$$

which give

$$i\psi_X + \psi_{TT} + \mu|\psi|^2\psi = 0, \tag{3.7}$$

$$\mu(\xi) = \frac{-g^3\Omega'\gamma^2\alpha_1}{\Omega^5\Omega''} \tag{3.8}$$

The dimensionless parameter μ is a monotonic increasing function of Kh, having the values zero and $(\Omega\gamma)^2$ for Kh = 1.363 and Kh $\to \infty$ respectively. The statement of the mathematical problem, given by Eq. (3.7), is completed by the following input condition at X = 0 (i.e. $x = x_\infty$: a reference point in infinitely deep water).

$$\psi(T,0) = 1 + 2\beta e^{i\alpha}\cos(2\pi T) \tag{3.9}$$

which corresponds to a system composed of a carrier-wave and a symmetric "side-band" disturbance

$$\zeta(t,x_\infty) = \frac{g\varepsilon}{\Omega^2} \text{Re}\{e^{-i\Omega t} + \beta e^{-i[(1+2\pi\varepsilon/\Omega\gamma)\Omega t - \alpha]} +$$

$$+ \beta e^{-i[(1-2\pi\varepsilon/\Omega\gamma)\Omega t - \alpha]}\} + O(\varepsilon^2) \tag{3.10}$$

For constant depth , μ = const, it is well-known that Eq. (3.7) with T in (0,1), subject to periodic boundary conditions has the following X invariants

$$J_1 = \int_0^1 |\psi|^2 dT \tag{3.11a}$$

$$J_2 = \int_0^1 (\psi^* \psi_T - \psi \psi_T^*) dT \tag{3.11b}$$

$$J_3 = \int_0^1 (|\psi|^4 - \frac{2}{\mu} |\psi_T|^2) dT \tag{3.11c}$$

These invariants are determined by the input condition, Eq. (3.9) so that

$$J_1 = 1 + 2\beta^2 \; ; \qquad J_2 = 0 \tag{3.12a,b}$$

$$J_3 = 1 + (4-P+2\cos 2\alpha) \cdot 2\beta^2 + 6\beta^4 \tag{3.12c}$$

where

$$P = 8\pi^2/\mu \tag{3.13}$$

For varying depth, $\mu = \mu(X)$, J_1 and J_2 remain invariant and are given by (3.11a,b) and (3.12a,b), but J_3 is a function of X governed by the equation,

$$\frac{dJ_3}{dX} = \frac{2\mu_X}{\mu^2} \int_0^1 |\psi_T|^2 dT \tag{3.14}$$

4. THREE-WAVES SYSTEMS

The solution of Eq. (3.7) can be expanded in a Fourier series

$$\psi(T,X) = \sum_{n=-\infty}^{\infty} D_n(X) e^{2\pi i n T} \qquad (4.1)$$

The boundary condition at $X = 0$, Eq. (3.9), gives $D_o(0) = 1$;
$D_1(0) = D_{-1}(0) = \beta e^{i\alpha}$; $D_n(0) = D_{-n}(0) = 0$ for $n \geq 2$.
Stiassnie and Kroszynski (1982) truncated the above given series and
considered only <u>three waves systems</u>:

$$\psi(T,X) = \sum_{n=-1}^{1} D_n(X) e^{2\pi i n T} \qquad (4.2)$$

Substituting Eq. (4.2) into Eq. (3.7) yields the following system of
<u>ordinary differential</u> equations:

$$i\frac{dD_o}{dX} + \mu[(|D_o|^2+4|D_1|^2)D_o+2D_1^2 D_o^*] = 0, \qquad (4.3a)$$

$$i\frac{dD_1}{dX} + \mu[(2|D_o|^2+3|D_1|^2-\frac{P}{2})D_1 + D_o^2 D_1^*] = 0. \qquad (4.3b)$$

Note that Eq. (3.11b) yields $D_{-1} = D_1$.

For constant depth the system of Eqs. (4.3a,b) has exact solutions
in terms of Jacobian elliptic functions with periods of order 1 in X which
is summarized in the Appendix; for details see Stiassnie and Kroszynski
(1982). These solutions depend on the invariants J_1,J_3, and on the parameter
μ, which in turn depends on the water depth h and on the modulation
period γ. For very mild depth variations, where $h_X = o(1)$, we apply

an asymptotic, WKB related approach, assuming the local solution to be
that of the constant depth type and using Eq. (3.14) to determine J_3.
J_1 and γ are fixed by the input conditions and J_3, is given through
$I(P)$ by:

$$\frac{dI}{dP} = -\sqrt{\frac{P(4-P)}{7}} \quad \frac{2 \ln\left(\dfrac{1 + \sqrt{\dfrac{4-P}{7P}}}{1 - \sqrt{\dfrac{4-P}{7P}}}\right)}{\ln \left(\dfrac{2P^2 (4-P)^2}{\left|(2P-1) I\right|}\right)} \tag{4.4}$$

where

$$I(P) = J_3(P) - J_1^2 \tag{4.5}$$

The initial value of I, at $X = 0$, where $P = P_o$ is denoted by I_o and
is given by:

$$I_o = \beta^2\left[2\beta^2 + 4(1+\cos2\alpha)-2P_o\right] \tag{4.6}$$

I_o, as well as $I(P)$ were assumed to be of $o(1)$ throughout the rather
lengthy derivation of Eq. (4.4). In all our examples we choose($\gamma = 2\pi\Omega^{-1}$)
$P_o = 2$, corresponding to the fastest growth-rate of the Benjamin-Feir
instability.

5. NUMERICAL VERIFICATION OF THE ASYMPTOTIC SOLUTION

In order to appraise the relevance of the asymptotic solution given in the previous section, we compare its results with those of a numerical solution of the system of ordinary differential equation (4.3a,b).

Fig. 5.1 shows $I=I(P)$ for four initial values of $I_0=I(2)$ ($I_0 = -0.04$, -0.01, 0.04, and 0.1). The broken line represents the asymptotic solution and was obtained by numerical integration of Eq. (4.4). The solid line was obtained, by substitution of the results obtained from a numerical solution of the system of O.D.E (4.3a.b) into the expression

$$I(P) = 2|D_1|^2\{|D_1|^2+2|D_0|^2-P+2|D_0|^2\cos[2(\arg D_1 - \arg D_0)]\} \quad (5.1)$$

Fig. 5.1 about here

The numerical solution of the system of O.D.E was obtained using a trapezoidal method and assumign the $P(h(X)) = 2+0.2X$. Note that the assumption $|I| \ll 1$, which is necessary for the asymptotic solution to be valid, imposes a restriction on the range of variation of P (P = 2.8 corresponds to $\Omega^2 h/g = 4$). In Fig. 5.2 we show three parts of the exterior group envelope $|\psi(0,X)|$ as well as the interior group envelope $|\psi(\frac{1}{2},X)|$ for the input conditions $\alpha = 0$, $\beta = 0.158$ ($I_0 = 0.1$).

Fig. 5.2 about here

Here again, solid lines represent the numerical solution of Eqs. (4.3a,b) with P = 2+0.2X while the broken lines correspond to results obtained by the asymptotic method, utilizing the relation:

$$|\psi(T,X)|^2 = \{-2I+2(4J_1-P)\tilde{z}-7\tilde{z}^2+ [S(2I+2P\tilde{z}-\tilde{z}^2)^{\frac{1}{2}}+4\tilde{z}\cos2\pi T]^2 \}/8\tilde{z} \quad (5.2)$$

where $\tilde{z}$ is given in the Appendix.

The three parts shown in Fig. 5.2 are for P=2, 2.44, 2.75 for the asymptotic solution compared to P in (2,2.03), (2.44, 2.49), (2.75, 281) for the numerical solution of the O.D.E.,respectively. The agreement between the two methods of solution, as seen in both the above figures is rather encouraging and seems to indicate the validity of our new asymptotic solution of the system (4.3a,b). Nevertheless, one still has to answer the question if, and to what extent, the syetem (4.3a,b) itself is a reasonable substitute for the N.L.S, Eq. (3.7) For constant depth Stiassnie and Kroszynski (1982) show a good quantitative agreement in the length of the modulation-demodulation cycle and only a qualitative agreement for the amplitudes. A similar trend can be seen in Fig. 5.3, which compares two numerical solutions, one for the N.L.S. (3.7) - dotted line, and the other for the system of O.D.E. (4.3a,b) - solid line.

Fig. 5.3 about here

The input data in Fig. 5.3 is $\alpha = 0$, $\beta = 0.1$ ($I_o = 0.04$) and the variation $P = 2 + 0.2X$ is assumed; Both the exterior and interior group envelopes are drawn.

We believe that our much-simplified asymptotic solution is not over-simplified, and is able to produce quite a few results of qualitative, and maybe even semi-quantitative relevance, which enable us some new physical insight.

6. ON P AND I_o -

One fundamental property of the asymptotic solution is that it depends on P and I_o solely. Given the input data α, β (and $P_o=2$), I_o is determined by Eq. (4.6). Then, integrating Eq. (4.4) from P_o to P the parameter I(P) is found, and the solution given by Stiassnie and Kroszynski (1982), is locally applied.

Fig. 6.1 about here

Fig. 6.1 gives the relation between P and nondimensional local water depth $K_\infty h$, (where $K_\infty = \Omega^2/g$ is the wave-number in infinitely deep water), for $\gamma = 2\pi\Omega^{-1}$. P is a monotonic decreasing function having the values infinity at $K_\infty h = 1.195$ (Kh=1.363) and 2 for $K_\infty h \to \infty$. The input value I_o, (for $P_o=2$) dependence on α and β is shown in Fig. 6.2. Note that different combinations of α and β give the same I_o, and thus basicly the same solution for any P.

Fig. 6.2 about here

7. GROUP ENVELOPES

The free-surface of an (unstable) shoaling wave-train, displays three distinct length scales: λ_1 – the <u>wave length</u>; λ_2 – the modulation or <u>group length</u>; and λ_3 – the modulation-demodulation, group-envelope, or maybe best '<u>supergroup' length.</u> These three lengths are given by

$$\lambda_1 = 2\pi/K, \tag{7.1a}$$

$$\lambda_2 = 2\pi\varepsilon^{-1}\Omega'/\Omega , \tag{7.1b}$$

$$\lambda_3 = \frac{4\varepsilon^{-2}P(\Omega')^3}{\Omega^2\Omega''\sqrt{P(4-P)}} \ln \left|\frac{2P^2(4-P)^2}{(2P-1)I}\right| , \tag{7.1c}$$

see Fig.(7.2a).

It can easily be seen that in the range of depths where the asymptotic solution applies, $K_\infty h \geqslant 4$, λ_1 and λ_2 remain almost constant. On the other hand, λ_3, which depends on P, exhibits quite a remarkable variation, as shown in Fig. 7.1.

<u>Fig. 7.1 about here</u>

In Fig. 7.1 we show the variation of λ_3 as a function of the depth $K_\infty h$ for four different input data I_o = -0.04, -0.01, 0.04, 0.1. For I_o <0, λ_3 decreases with decreasing depth, but for $I_o > 0$ λ_3 increases with decreasing depth up to a "critical depth" (corresponding to I = 0) and from there on starts to decrease.

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Fig. 7.2 shows the group envelopes (dashed line) and wave
envelope (solid lines) at a fixed instant for $\epsilon = 0.2$, at the following
four locations: (a) - infinitely deep water, $P_o = 2$, $I_o = 0.1$;
(b) $K_\infty h = 11.2$, $P = 2.2$, $I = 0.052$; (c) $K_\infty h = 5.7$, $P = 2.45$, $I = 0.01$;
(d) $K_\infty h = 4.2$, $P = 2.75$, $I = -0.028$.

Fig. 7.2 about here

In Fig. 7.2a we have added a portion of the wavy-surface (thin solid
line) as well as the lengths λ_1, λ_2, and λ_3. Note that the supergroups
(namely: the exterior and interior group envelopes) are fixed in
space, while the wave envelope moves with group-velocity and the waves
themselves with the phase velocity. Similar sketches to Fig. 7.2
were obtained for the three other cases given in Fig. 7.1.

In order to complete the picture for shallower water depth we
present in Fig. 7.3 the group envelopes at five locations: $P = a:(2, 2.05)$,
b:(2.32, 2.63), c:(17.7, 22.8), d:(-22.4,-17.5), e:(-4.1,-3.9),as
obtained from a numerical solution of the system (4.3a,b) for the
same input data as in Fig. 7.2 assuming $\mu = 39.47-10X$.

Fig. 7.3 about here

The results in Fig. 7.3 indicate that λ_3 continues to shorten
and that the intensity of modulation decreases.

8. THE MEAN FLOW FIELD

We express the mean flow $u = \partial\phi_o/\partial x = -\partial\tilde{\psi}/\partial z$, $v = \partial\phi_o/\partial z = \partial\tilde{\psi}/\partial x$

through the stream function $\tilde{\psi}$

$$\tilde{\psi} = \frac{g^2 K\gamma\epsilon^{-1}}{4\pi\Omega^3 h} \left(|D_o|^2 + 2|D_1|^2 \right)Z +$$

$$+ \frac{g^3\gamma\epsilon^{-1}}{4\pi\Omega^3\Omega'} \cdot \frac{\beta_2 (D_o^* D_1 + D_o D_1^*) \, \text{sh}(2(Z+H))\cos(2\pi T)}{\frac{g\gamma}{2\pi\epsilon\Omega'}\,\text{sh}(2H) - \text{ch}(2H)} +$$

$$+ \frac{g^2\gamma\epsilon^{-1}}{4\pi\Omega^3\Omega'} \cdot \frac{\beta_2 |D_1|^2 \text{sh}(4(z+H))\cos(4\pi T)}{\frac{g\gamma}{2\pi\epsilon\Omega'}\,\text{sh}(4H) - 2\text{ch}(4H)} + \text{constant} \qquad (8.1)$$

where $Z = \pi\epsilon z/\Omega'\gamma$, $H = \pi\epsilon h/\Omega'\gamma$ and $\tilde{\psi} = \epsilon^{-2}\Omega^3\tilde{\psi}/g^2$ are dimensionless

quantities. The constant in Eq. (8.1) is chosen so that $\tilde{\psi} = 0$

at the bottom. The mean free surface ζ_o is given by Eq. (2.11)

Figures 8.1, 8.2, 8.3 about here

The stream-function $\Psi(T,Z)$ as well as the mean free-surface

for cases a, b and d of Fig. 7.2 are presented in Figures 8.1, 8.2

and 8.3 respectively. These figures demonstrate the rather complicated

structure of the wave induced mean flow field.

Some of the main features are: (i) the mean current, which is

shown in part (b) of the figures, as well as the mean free surface,

in part (c) exhibit a somewhat cellular structure influenced by the

wave envelope variations; (ii) a dominant adverse cu rent appears

underneath the high waves and a much weaker, positive current (in the

wave propagation direction) under the low waves, for the shallower cases

the positive currents almost disappears; (iii) the magnitude of the

maximum adverse currents at the free surface is almost the same for

all three depths ($K_\infty h = \infty$, 11.2 and 4.2); (iv) one can notice the tendency of the flow fields to become more uniform in their lower parts and on the sides of the supergroups (where the modulation amplitudes get much smaller); (v) there is a set-down in the mean free surface accompanying the peaks of the wave envelope and a smaller set-up accompanying their troughs.

APPENDIX - Solution of Eqs. (4.3a,b) for constant depth

Input data: $\quad J_1, J_3, P; \qquad\qquad I = J_3 - J_1^2$

$$C_1 = \frac{1}{7}(4J_1-P)\left\{1-\left[1-\frac{14I}{(4J_1-P)^2}\right]^{\frac{1}{2}}\right\} \qquad (A.1)$$

$$C_2 = P \cdot \left[1-(1+\frac{2I}{P^2})^{\frac{1}{2}}\right] \qquad (A.2)$$

$$c = \max(C_1,C_2); \qquad\qquad d = \min(C_1,C_2) \qquad (A.3a,b)$$

$$a = 2P; \qquad\qquad b = \frac{2}{7}(4-P) \qquad (A.3c,d)$$

$$k^2 = 1 - \frac{a-b}{a}\cdot\frac{c-d}{b} \qquad (A.4)$$

$$y = -\frac{2\pi^2}{P}\sqrt{7ab}\cdot X \qquad (A.5)$$

$cd(y,k)$ is a Jacobian elliptic function of the argument y with modulus k.

$$\tilde{z} = \frac{a\cdot b\cdot(1-cd^2)}{a-b\cdot cd^2} \qquad (A.6)$$

$$|D_1| = \sqrt{\tilde{z}/2}; \qquad\qquad |D_o| = \sqrt{J_1-\tilde{z}} \qquad (A.7a,b)$$

$$\cos[2(\arg D_1-\arg D_o)] = \frac{I-1.5\tilde{z}^2+(P-2J_1)\cdot\tilde{z}}{2\tilde{z}(J_1-\tilde{z})} \qquad (A.7c,d)$$

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List of Figures

Figure 5.1

Figure 5.2

Figure 5.3

$K_\infty h = 1.195$

$K_\infty h$

P

Figure 6.1

Figure 6.2

$$\frac{\varepsilon^2 \lambda_3 K_\infty}{}$$

$K_\infty h$

Figure 7.1

$I_o = 0.1$

$I_o = -0.04$

$I_o = -0.01$

$I_o = 0.04$
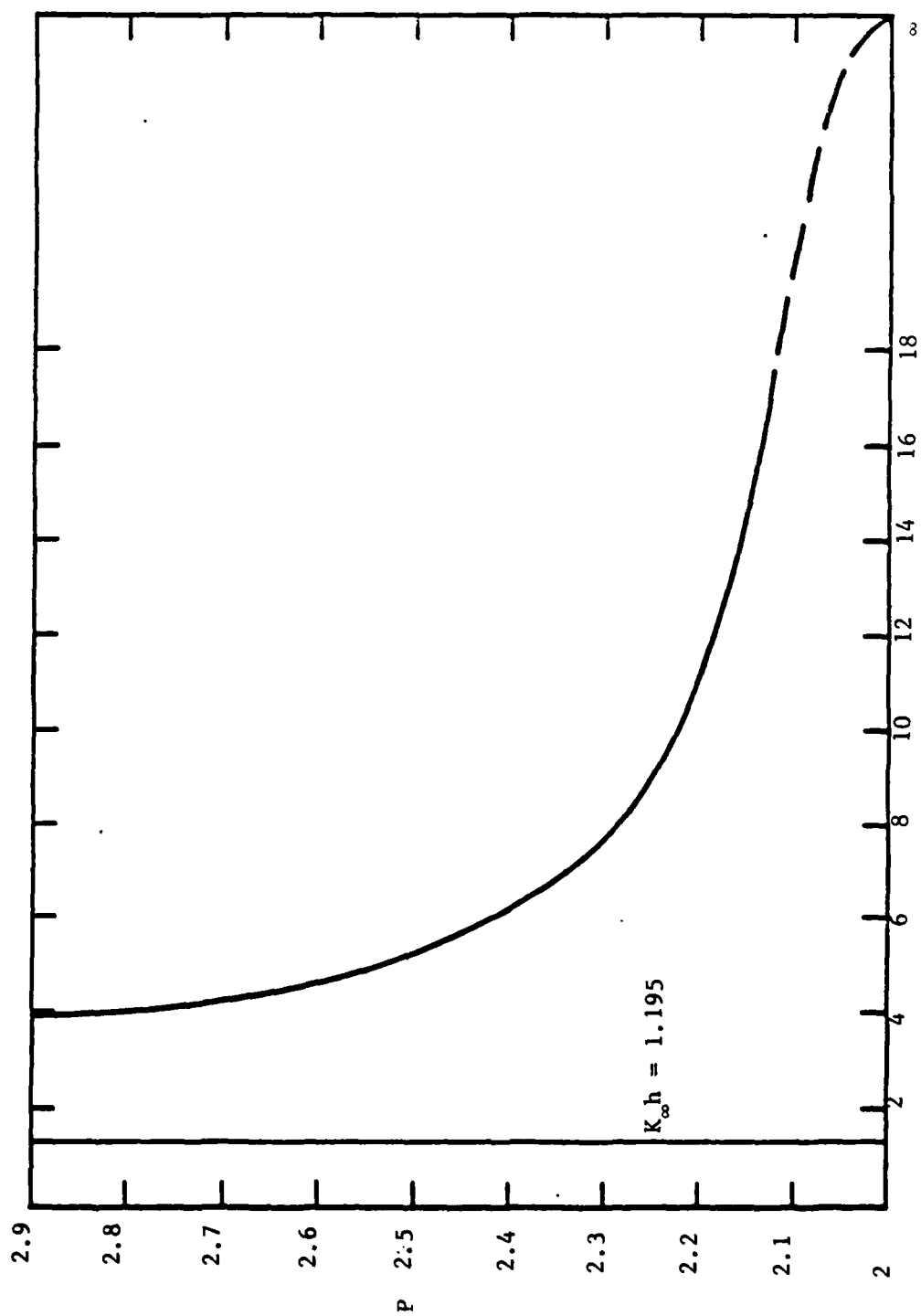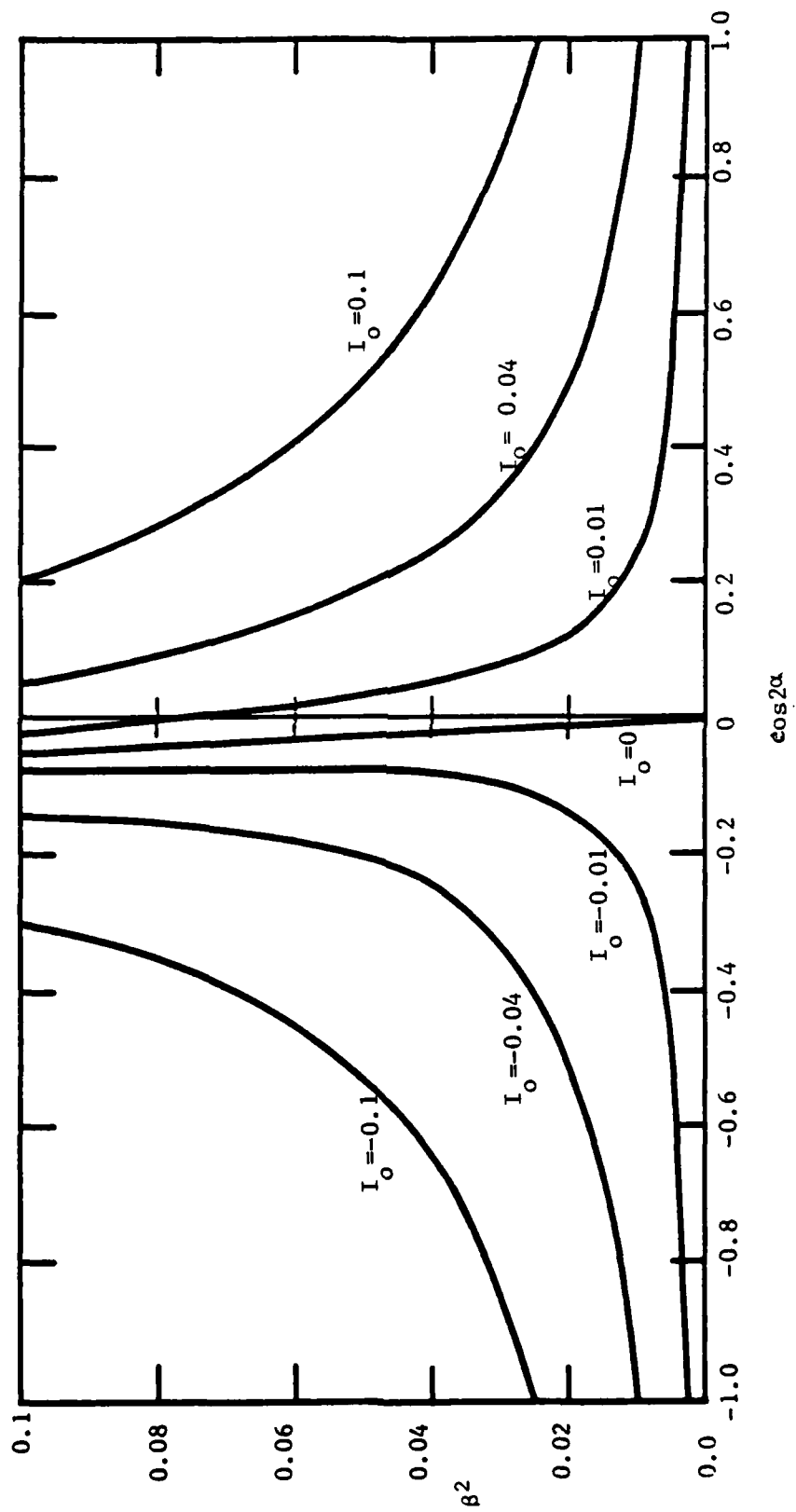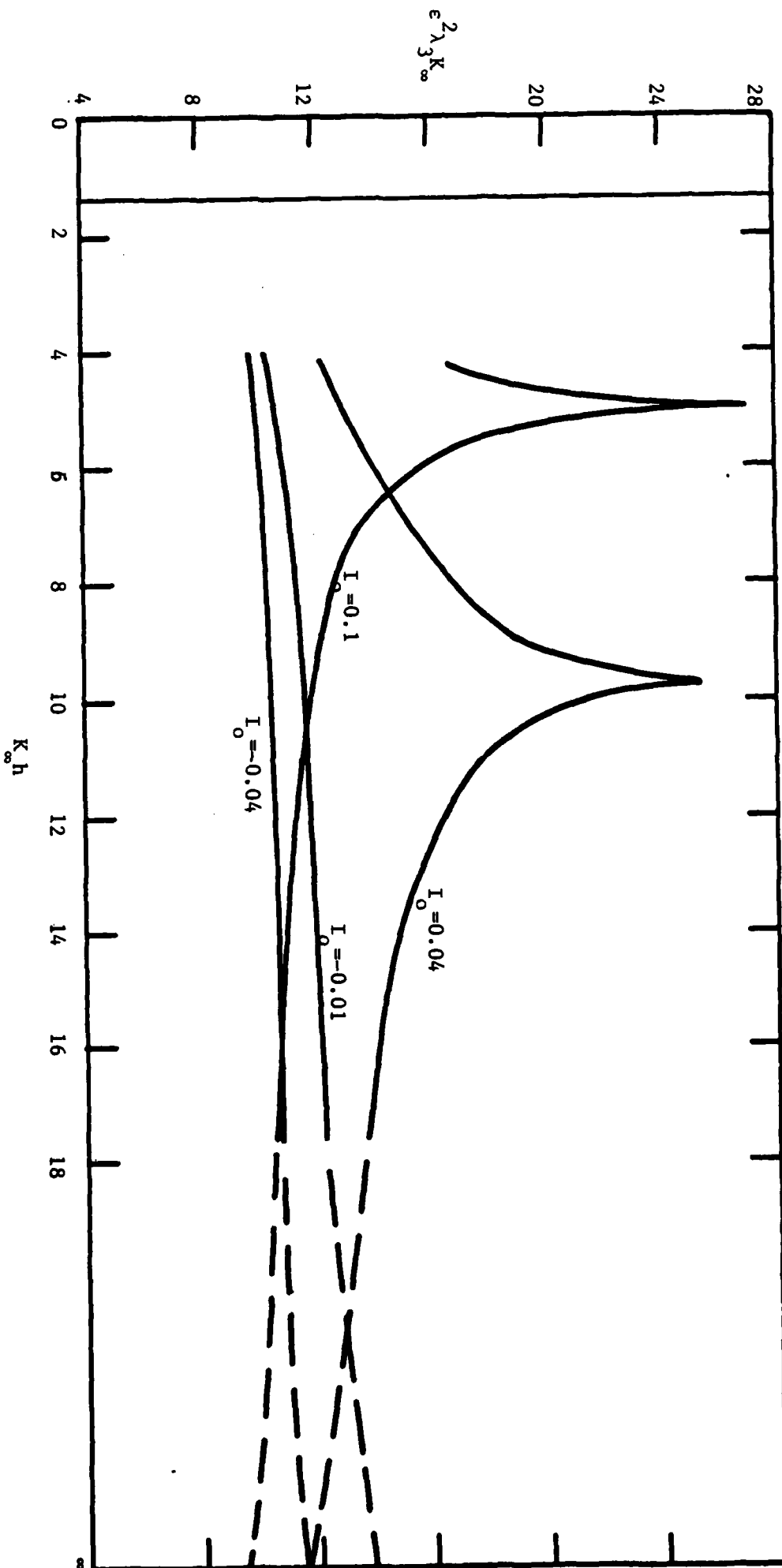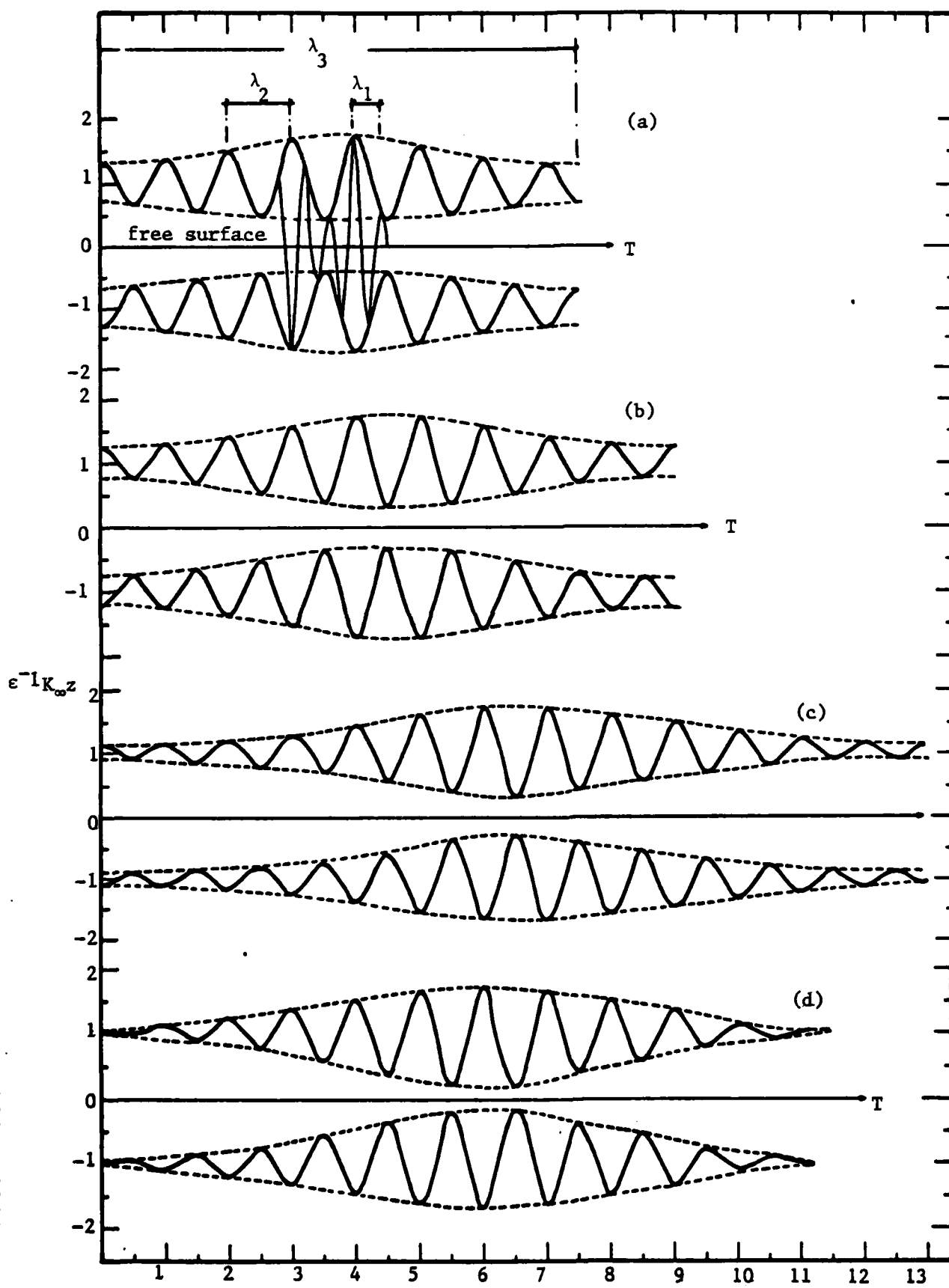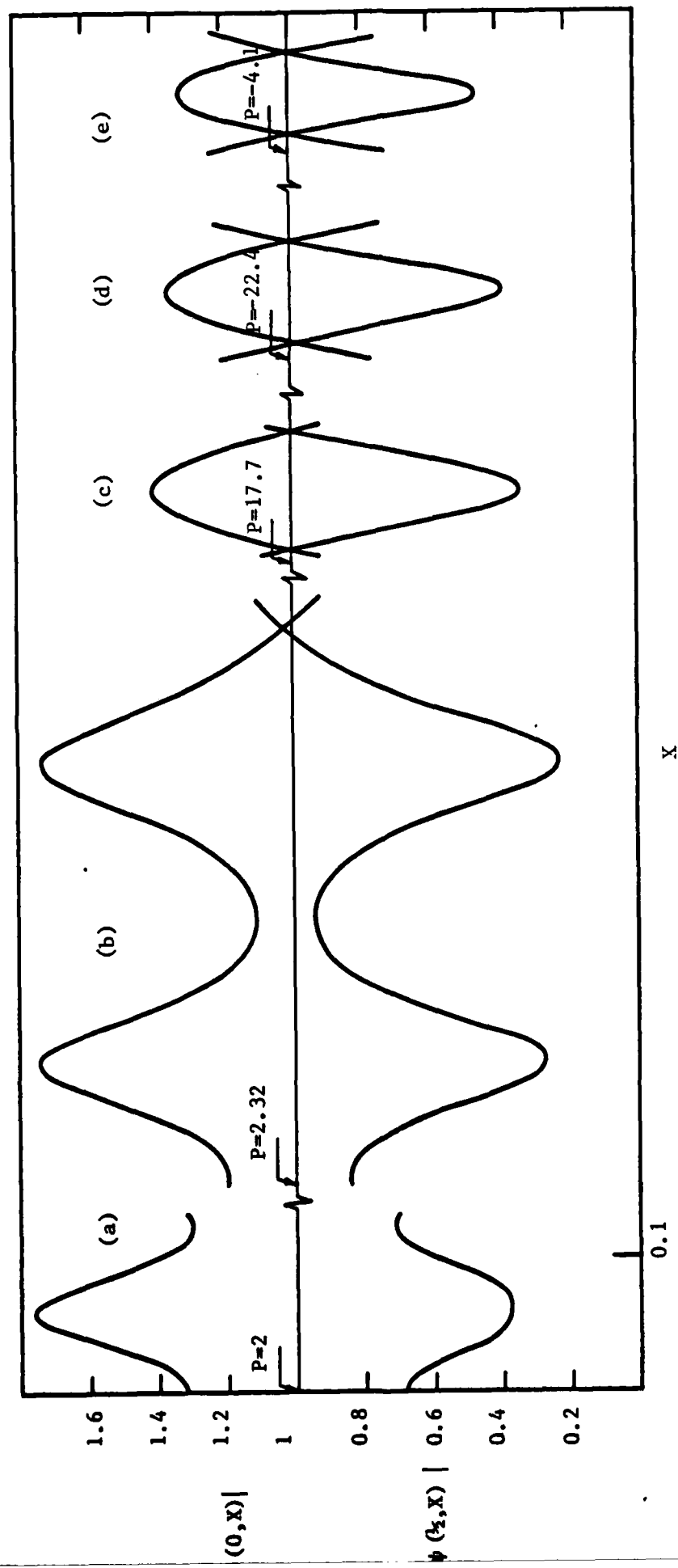
$$T = \frac{\varepsilon K_\infty}{\pi} (x - x_\infty)$$

Figure 7.2

Figure 7.3

Figure 8.1
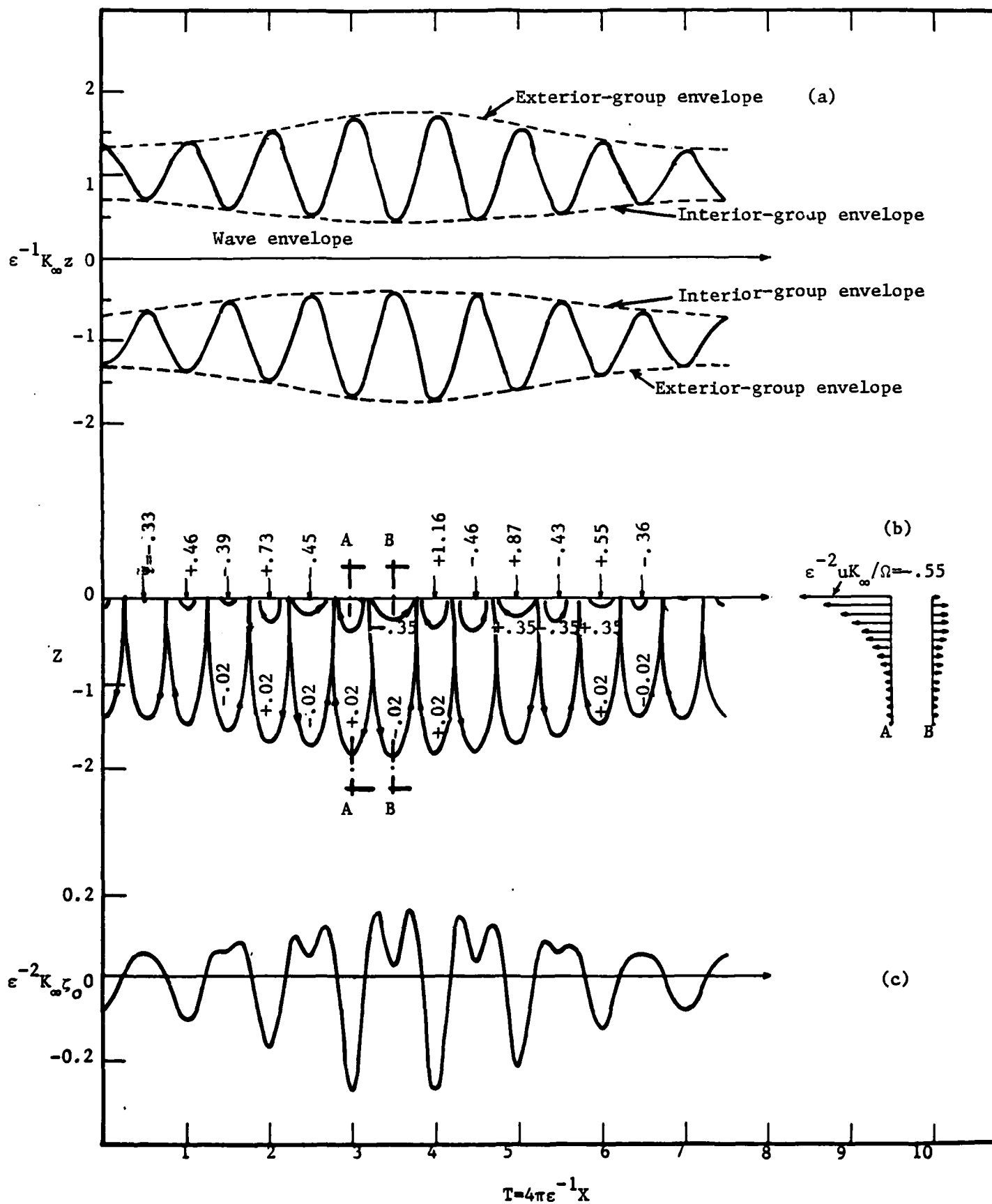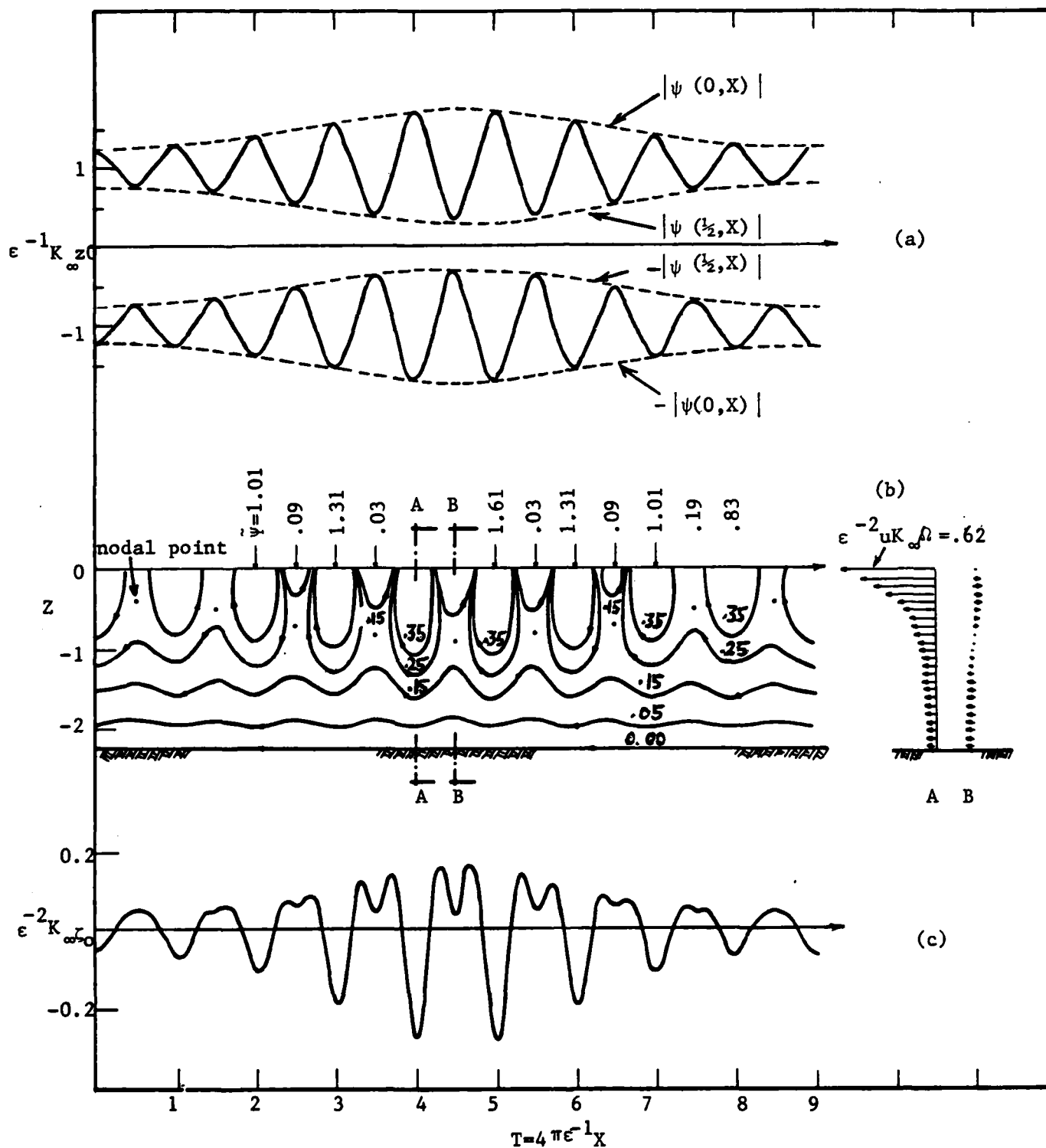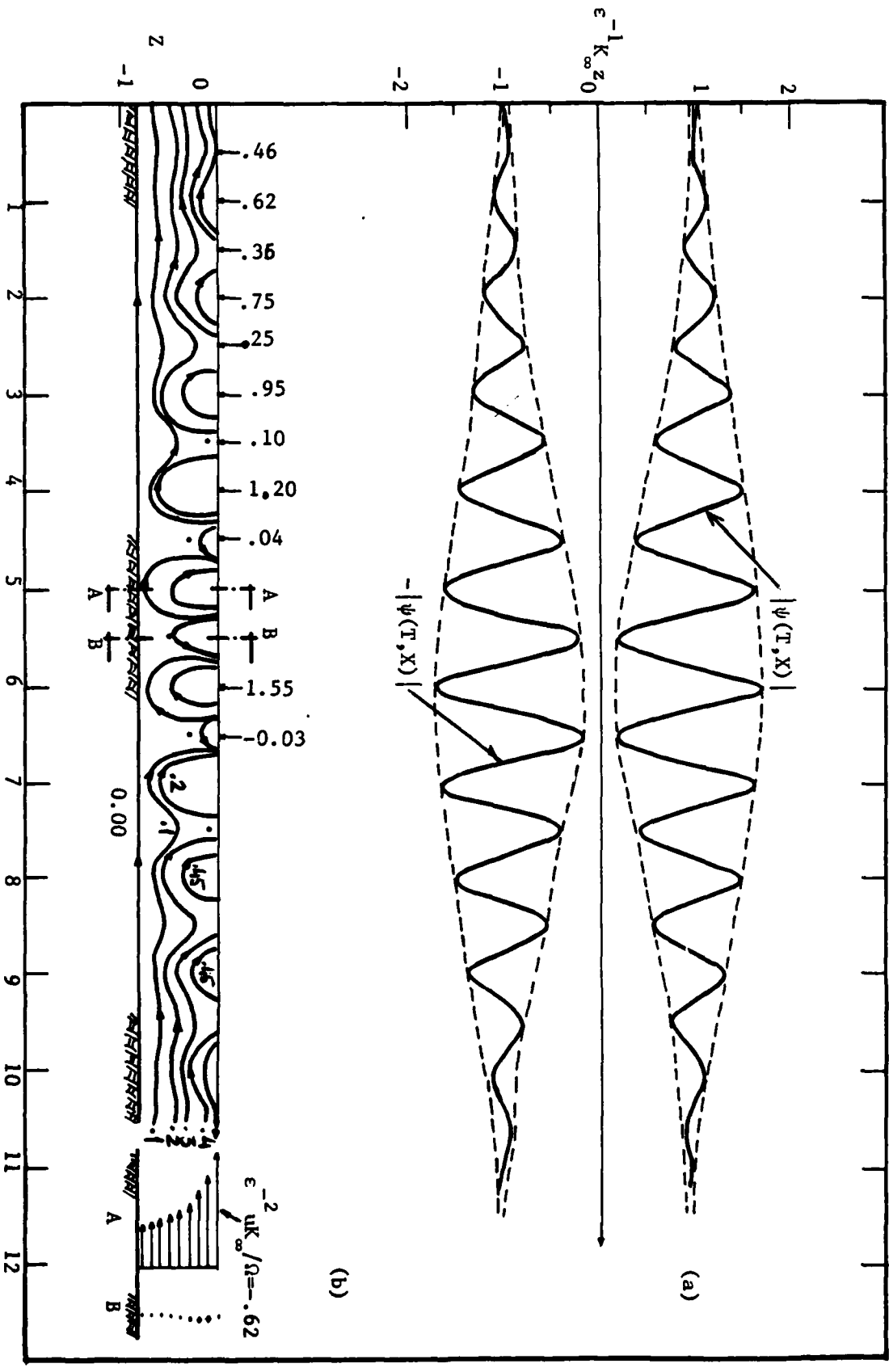
Figure 8.2

$$T = 4\pi\epsilon^{-1}$$

Figure 8.3